

GTWS meeting 2022

May 17, 2022(Tue) 12:20~12:40

Air-sea gas exchange in a seagrass ecosystem



Photo from NOAA's HP

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Seagrass's role in Global carbon cycle

	Burial Tg C y ⁻¹
Vegetated habitats	
Mangroves	23.6
Salt Marsh	60.4
Seagrass	27.4
Total vegetated habitats	111.4
Depositional areas	98.9
Total coastal burial	210.3
Deep sea burial	6.0
Total oceanic burial	216.3

(Duarte et al., 2005)

Seagrass buries 27.4 Tg C year⁻¹, ~10 % of total ocean burial

Calculation of CO₂ flux across air and sea

$$\text{CO}_2 \text{ flux} = k \cdot ([CO_{2\text{sea}}] - [CO_{2\text{air}}])$$

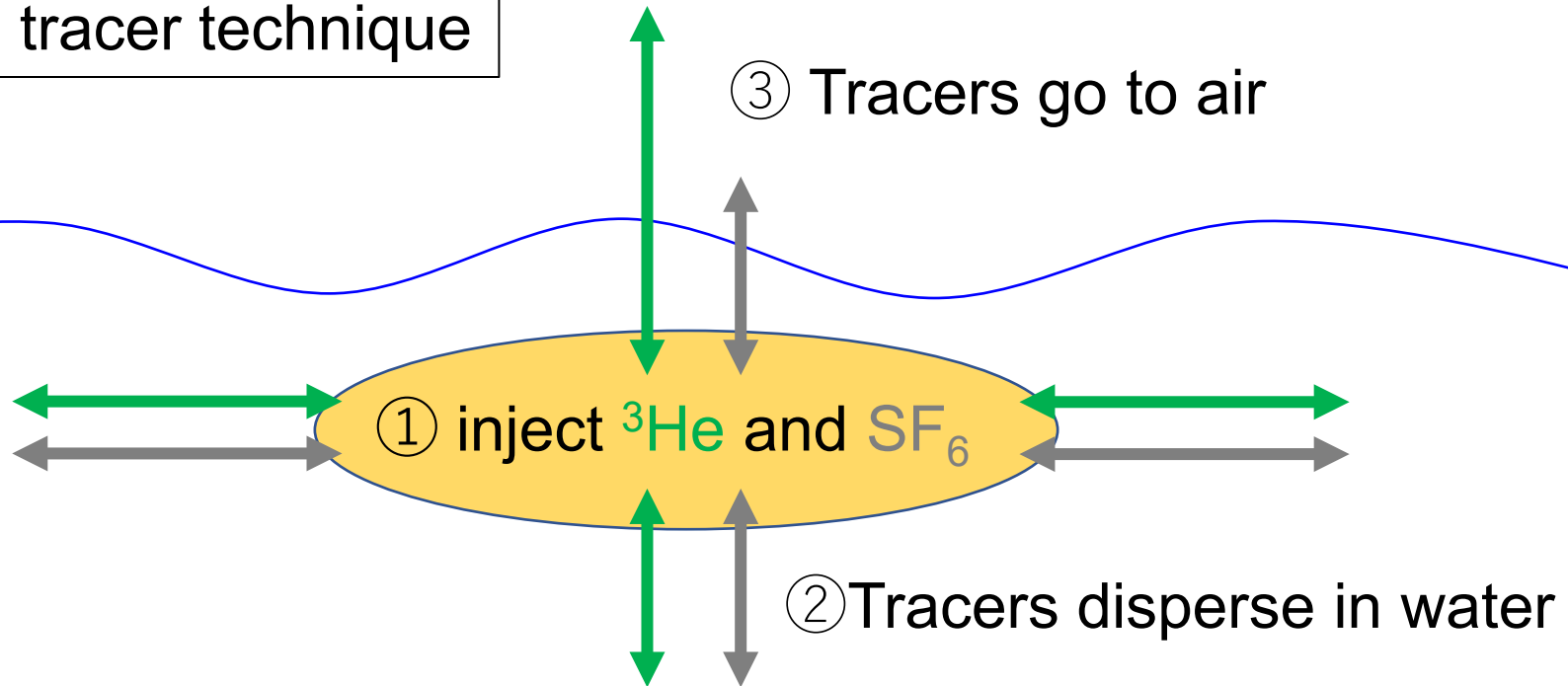
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gas transfer velocity

CO₂ conc.

- Gas transfer velocity (k) is hard to measure.
- In offshore region, k well parameterized by wind (u)
 $k = 0.251 \times u_{10m}^2$ for CO₂ at 20 °C (e.g., Wanninkhof, 2014)
- In shallow region, other parameters, such as current speed, important as well (e.g., Ho et al., 2016).

Methods How measure the gas transfer velocity (k)?

Dual tracer technique

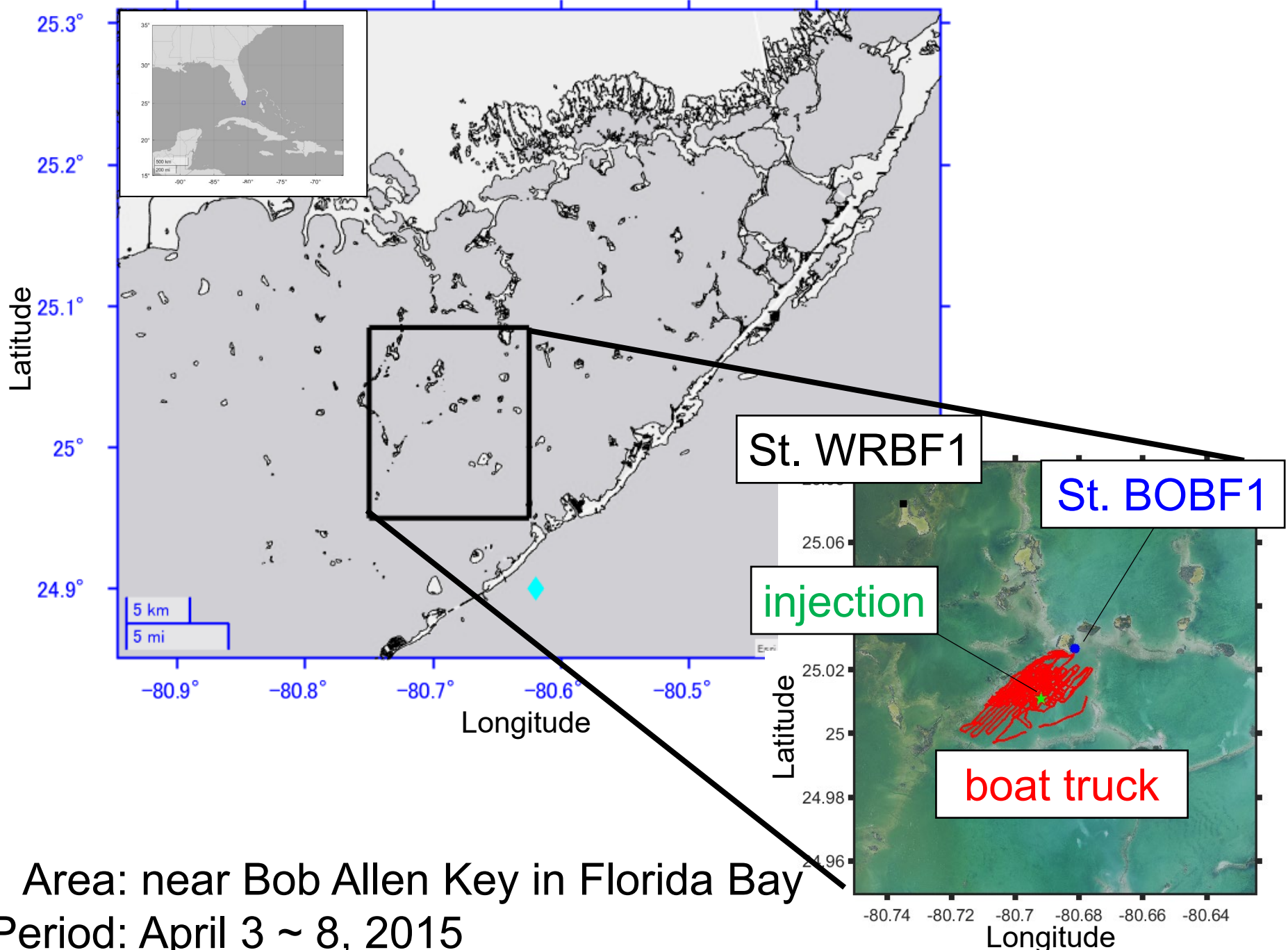


$$\underline{k(600)} = -h \frac{d}{dt} \left(\ln({}^3\text{He}_{exc} / \text{SF}_6) / 1 - \left(Sc_{\text{SF}_6} / Sc_{{}^3\text{He}} \right)^{-\frac{1}{2}} \right) \left(600 / Sc_{{}^3\text{He}} \right)^{-\frac{1}{2}}$$

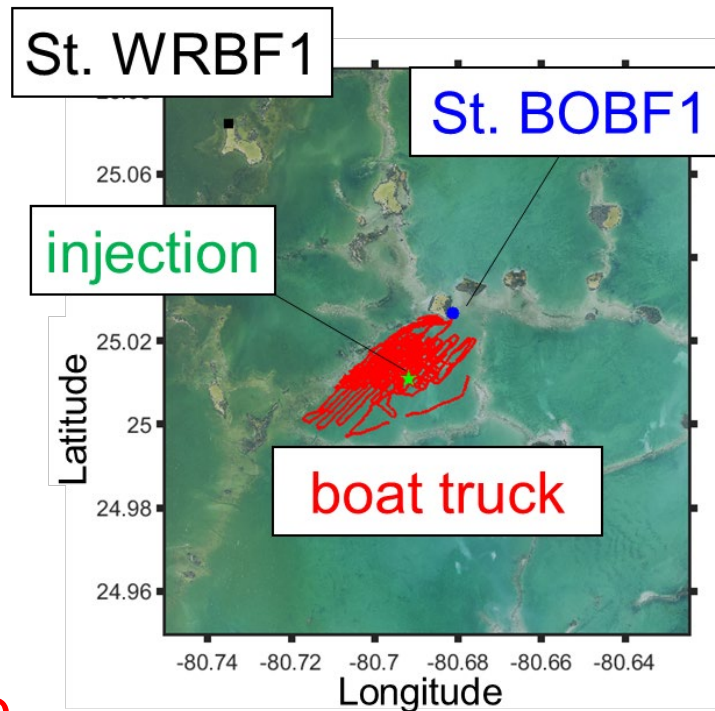


k for CO_2 at 20 °C, fresh water

Methods



Methods

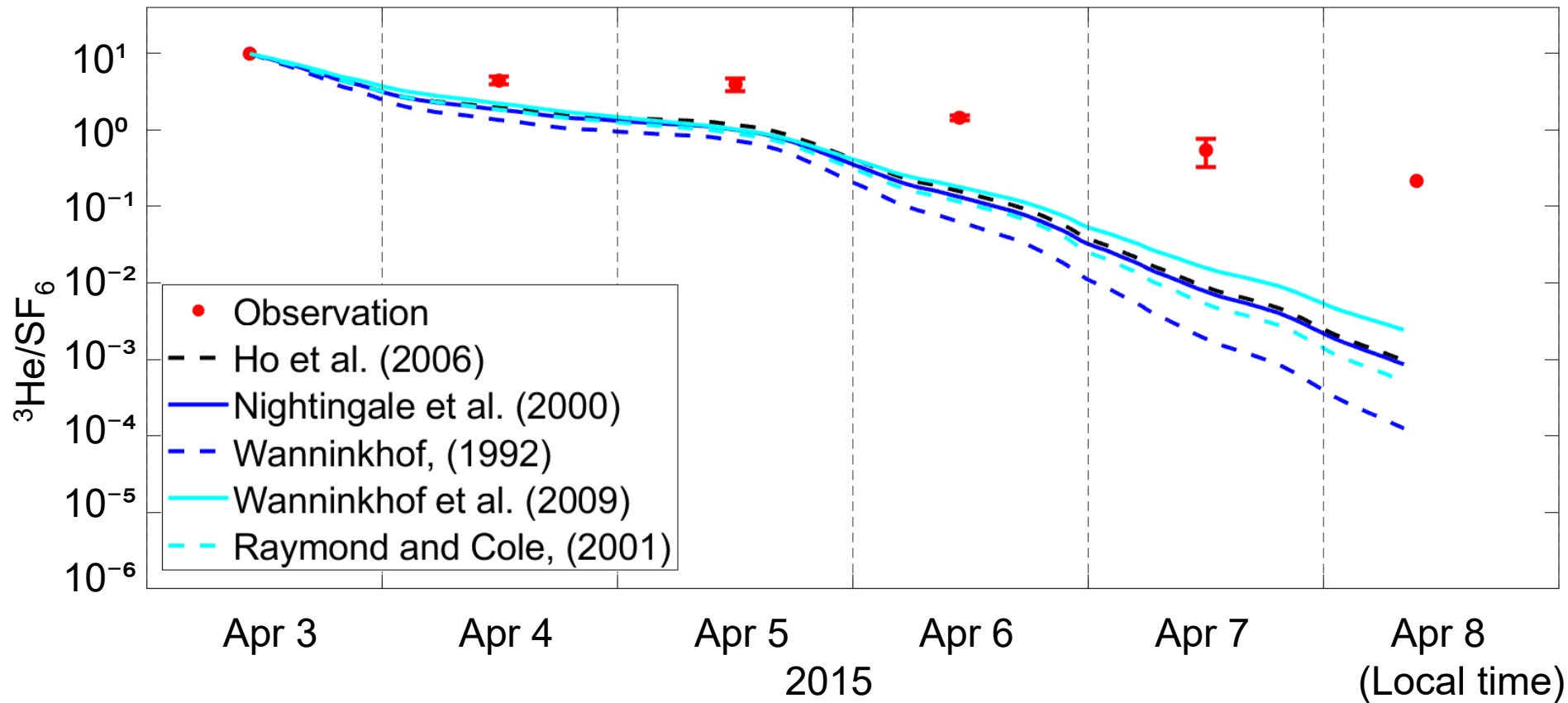


boat → water sample

parameters	method
^3He conc.	He isotope mass spectrometer (Ludin et al., 1998)
SF_6 conc.	gas chromatograph equipped with an electron capture detector (GC/ECD) (Wanninkhof et al., 1987)
Air, water pCO_2	non-dispersive infrared (NDIR; LI-COR 840A) analyzer (Ho et al., 1997; Pierrot et al., 2009)

Result

change in ${}^3\text{He}/\text{SF}_6$ with time



$$k(600) = -h \frac{d}{dt} \left(\ln({}^3\text{He}_{exc}/\text{SF}_6) / 1 - \left(S_{C_{\text{SF}_6}} / S_{C_{{}^3\text{He}}} \right)^{-\frac{1}{2}} \right) \left(600 / S_{C_{{}^3\text{He}}} \right)^{-\frac{1}{2}}$$

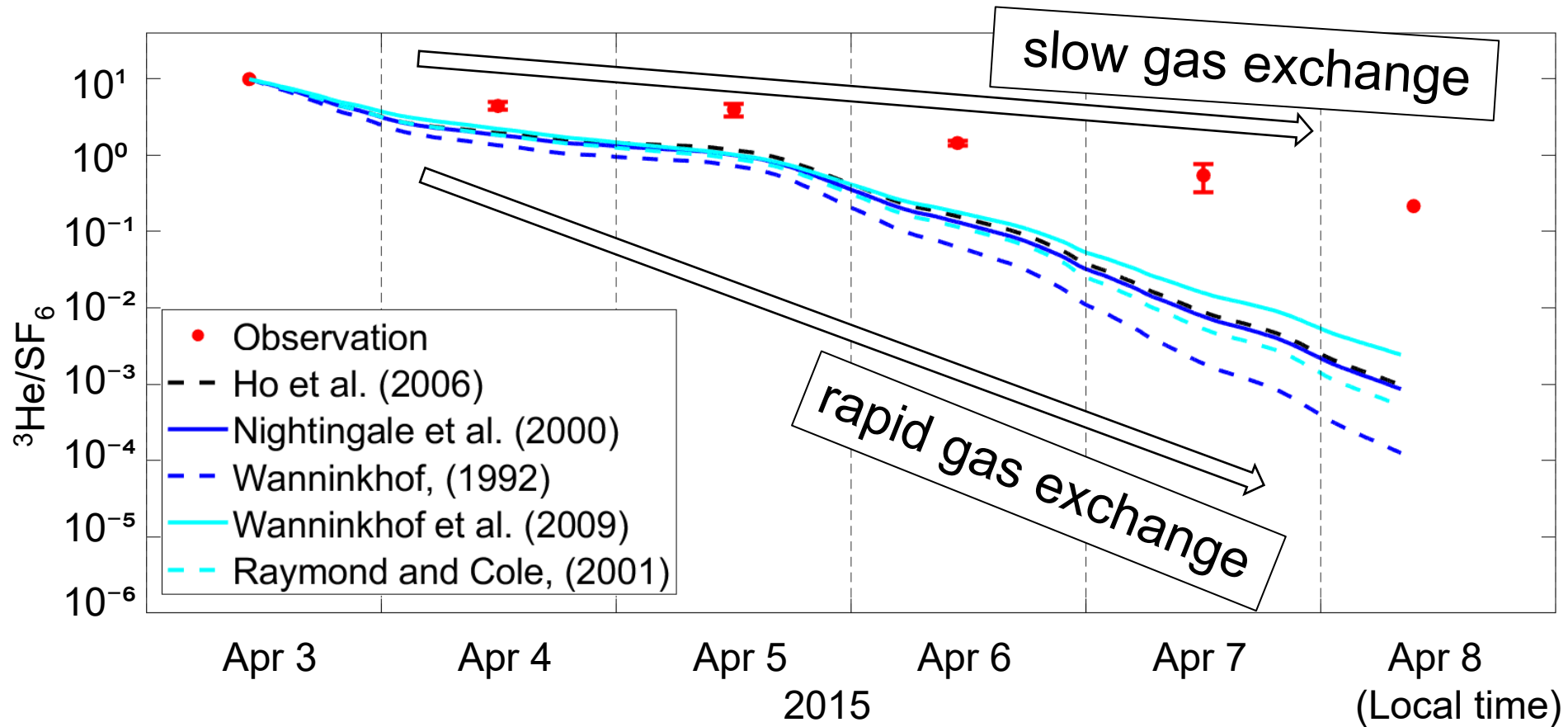
$k(600) = 4.7 \pm 1.8 \text{ cm hr}^{-1}$, lower than other coastal & offshore at same u_{10m}

(Ho and Wanninkhof, 2016)

↑
 k for CO_2 at 20°C , fresh water

Result

change in $^3\text{He}/\text{SF}_6$ with time



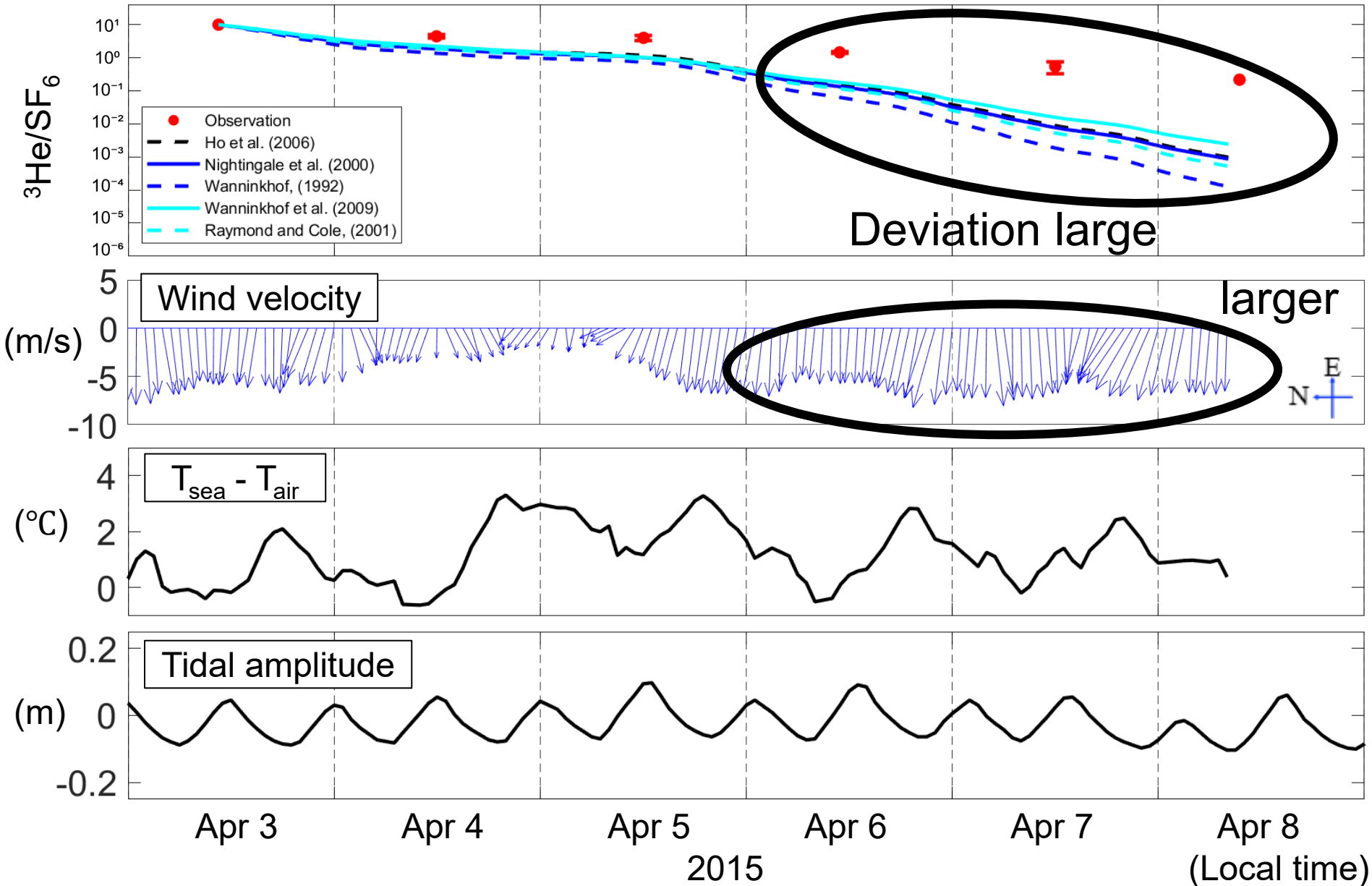
$$k(600) = -h \frac{d}{dt} \left(\ln(^3\text{He}_{exc}/\text{SF}_6) / 1 - \left(Sc_{\text{SF}_6} / Sc_{^3\text{He}} \right)^{-\frac{1}{2}} \right) \left(600 / Sc_{^3\text{He}} \right)^{-\frac{1}{2}}$$

Published parameterizations decrease rapidly than observation



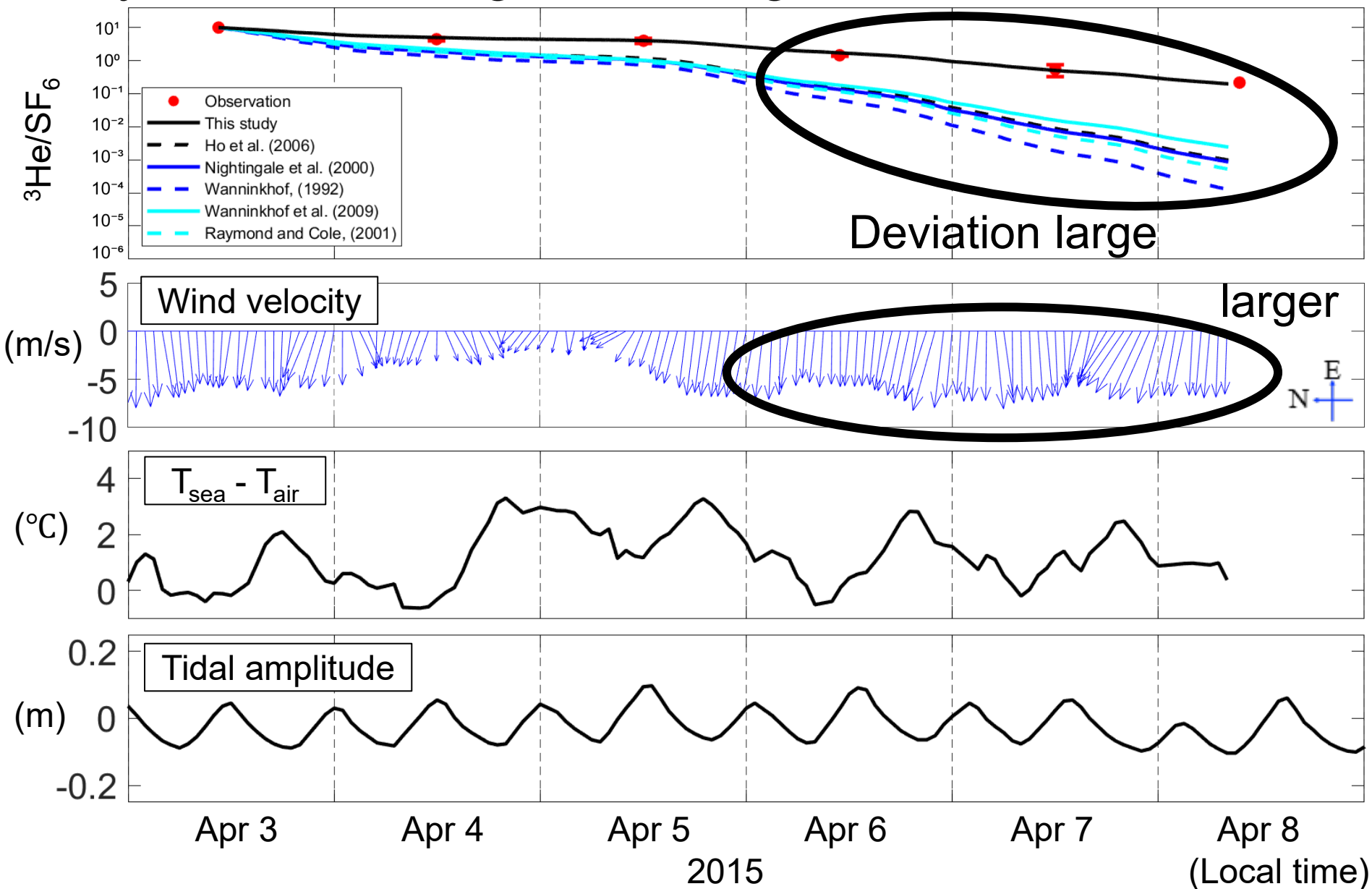
They overestimate gas transfer velocity (k)

Why lower air-sea gas exchange ?



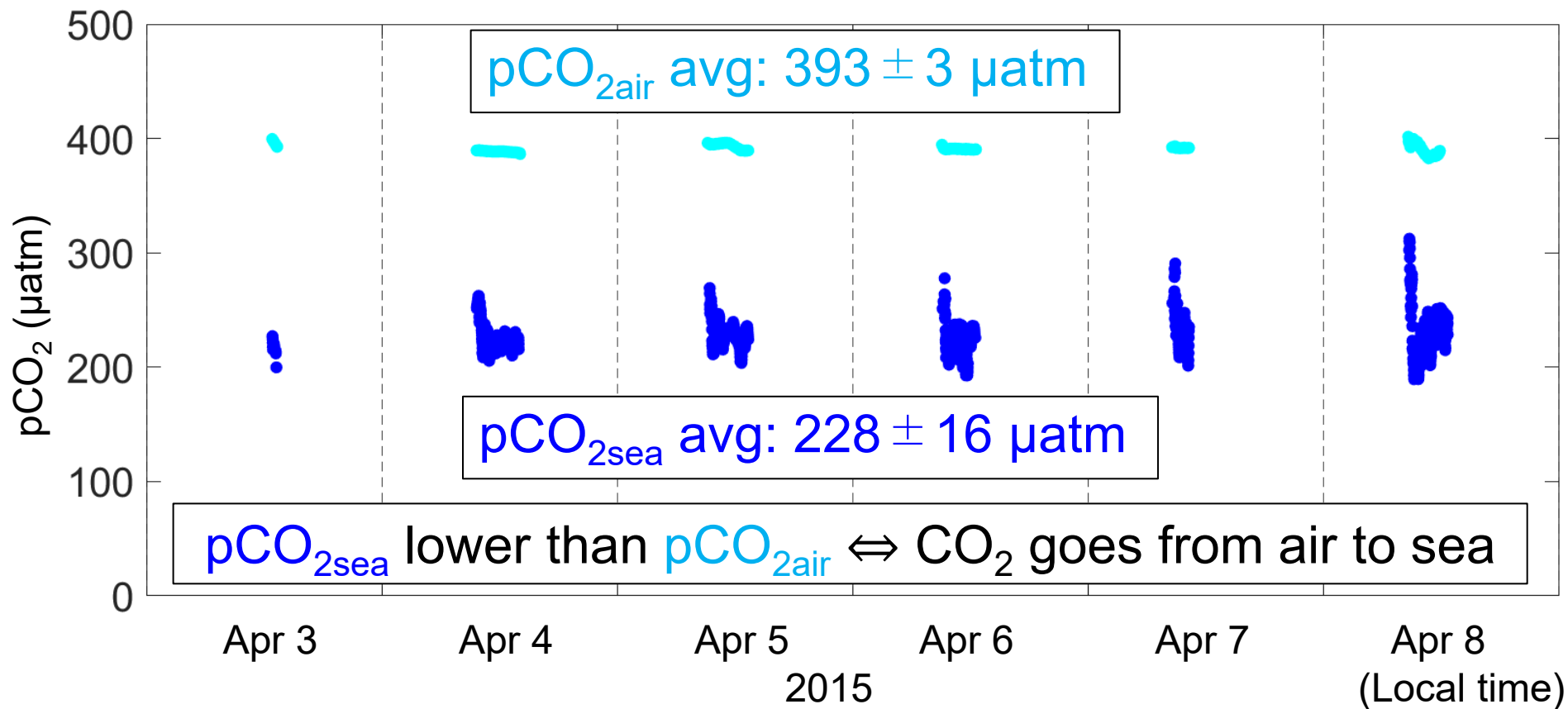
- Wind is dominant factor
- Lower $k \Leftarrow$ fetch limitation or wave attenuation by seagrasses

Why lower air-sea gas exchange ?



New parameterization, $k(600) = 0.122 \cdot u_{10}^2$ fits well.

Measured pCO₂ and CO₂ flux



Values are consistent with Zhang et al. 2014

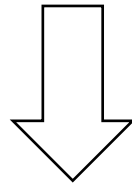
$$\text{CO}_2 \text{ flux} = k \cdot ([\text{CO}_{2\text{sea}}] - [\text{CO}_{2\text{air}}]) = -4.3 \pm 2.6 \text{ mmol m}^{-2} \text{ day}^{-1}$$

Annual CO₂ flux

Zhang et al. 2014 measured pCO₂

They applied $k(660) = 0.31 \cdot u_{10}^2$ (Wanninkhof, 1992)

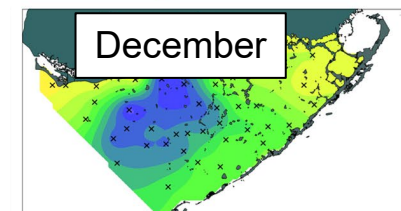
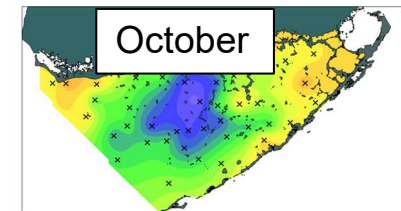
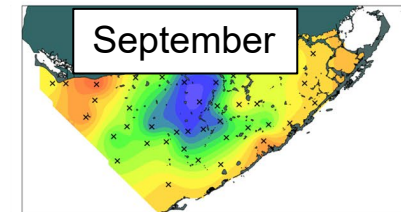
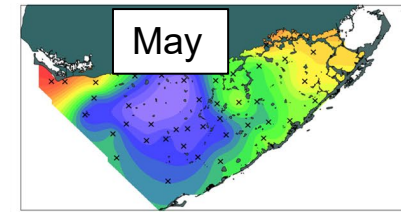
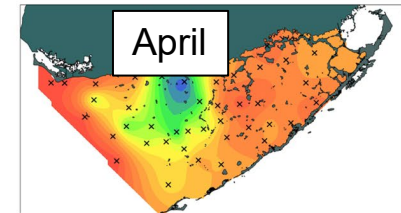
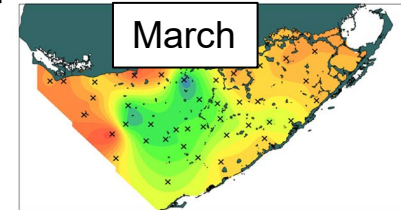
$$\text{Annual CO}_2 \text{ flux} = \mathbf{3,930} \pm 910 \text{ mmol m}^{-2} \text{ yr}^{-1}$$



Recalculate by $k(600) = 0.122 \cdot u_{10}^2$

$$\text{Annual CO}_2 \text{ flux} = \mathbf{1,500} \pm 350 \text{ mmol m}^{-2} \text{ yr}^{-1}$$

$$\mathbf{1,500 / 3,930 = 38\%}$$



(Zhang et al., 2014) pCO_{2,sea} (µatm)

Summary

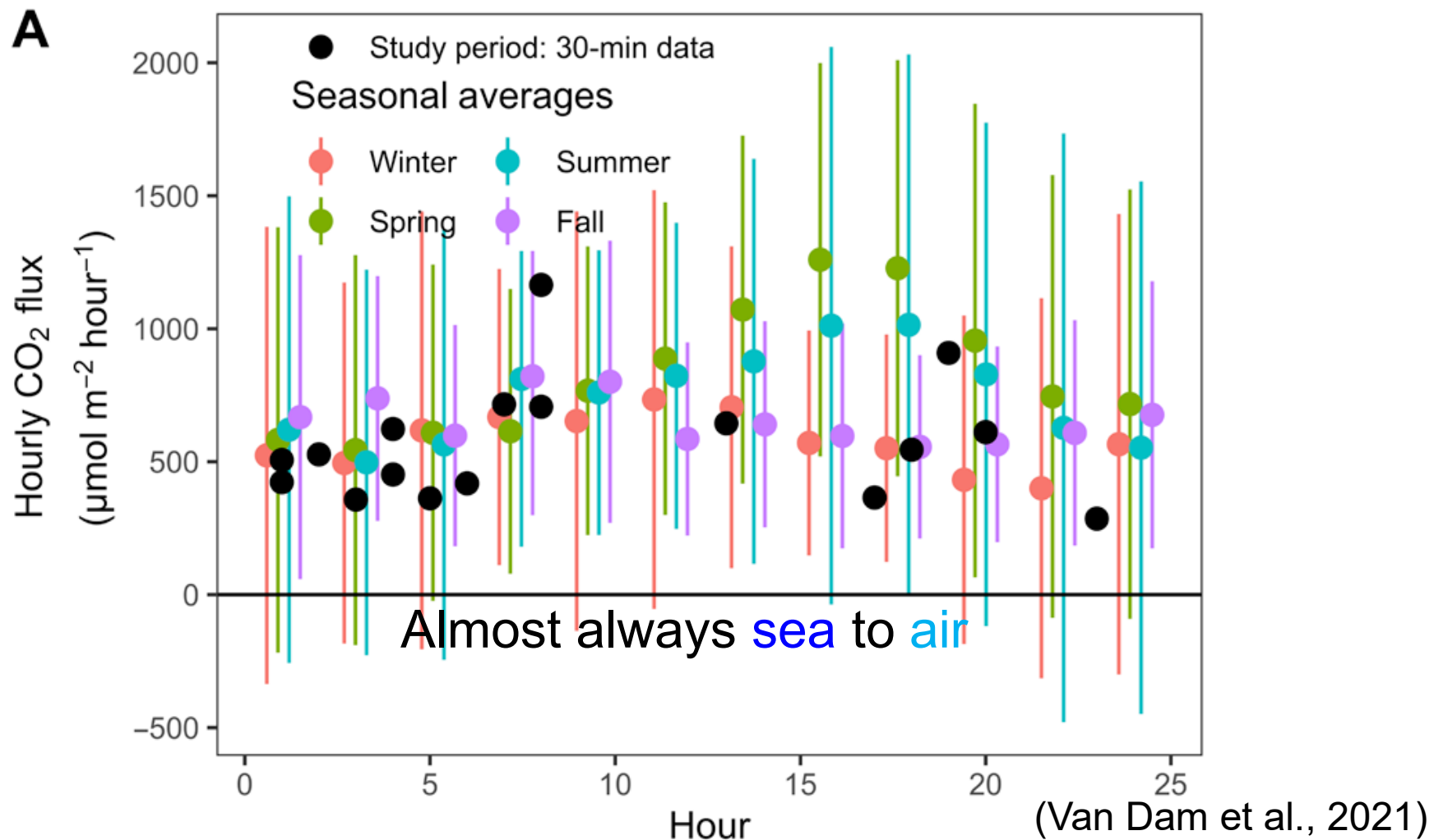
- Gas transfer velocity (k) lower than other coastal & open oceans at same wind speed
- Specific parameterization for k is needed for inland ecosystems
- New parameterization, $k(600) = 0.122u_{10}^2$, fitted air-sea gas exchange well.
- CO₂ flux during observation = -4.3 ± 2.6 mmol m⁻² day⁻¹
Recalculated annual CO₂ flux = $1,500 \pm 35$ mmol m⁻² yr⁻¹

$$\text{cvRMSE} = \frac{\sqrt{\frac{1}{N} \sum_{n=1}^N (R_{\text{mod}}^n - R_{\text{obs}}^n)^2}}{\overline{R_{\text{obs}}}}$$

References	Parameterization	Mean k(600) (cm h ⁻¹)	cvRMSE
This study	$k(600) = 0.122u_{10}^2$	5.4 ± 2.9	8.6%
Ho et al. (2006)	$k(600) = 0.266u_{10}^2$	11.7 ± 6.4	84.4%
Nightingale et al. (2000)	$k(600) = 0.333u_{10} + 0.222u_{10}^2$	11.8 ± 6.0	88.4%
Wanninkhof (1992)	$k(660) = 0.31u_{10}^2$	14.2 ± 7.8	99.4%
Wanninkhof et al. (2009)	$k(660) = 3 + 0.1u_{10} + 0.064u_{10}^2 + 0.011u_{10}^3$	10.5 ± 4.5	83.1%
Raymond and Cole (2001)	$k(600) = 1.58e^{0.3u_{10}}$	12.4 ± 6.6	90.0%

Compare CO₂ flux with previous study

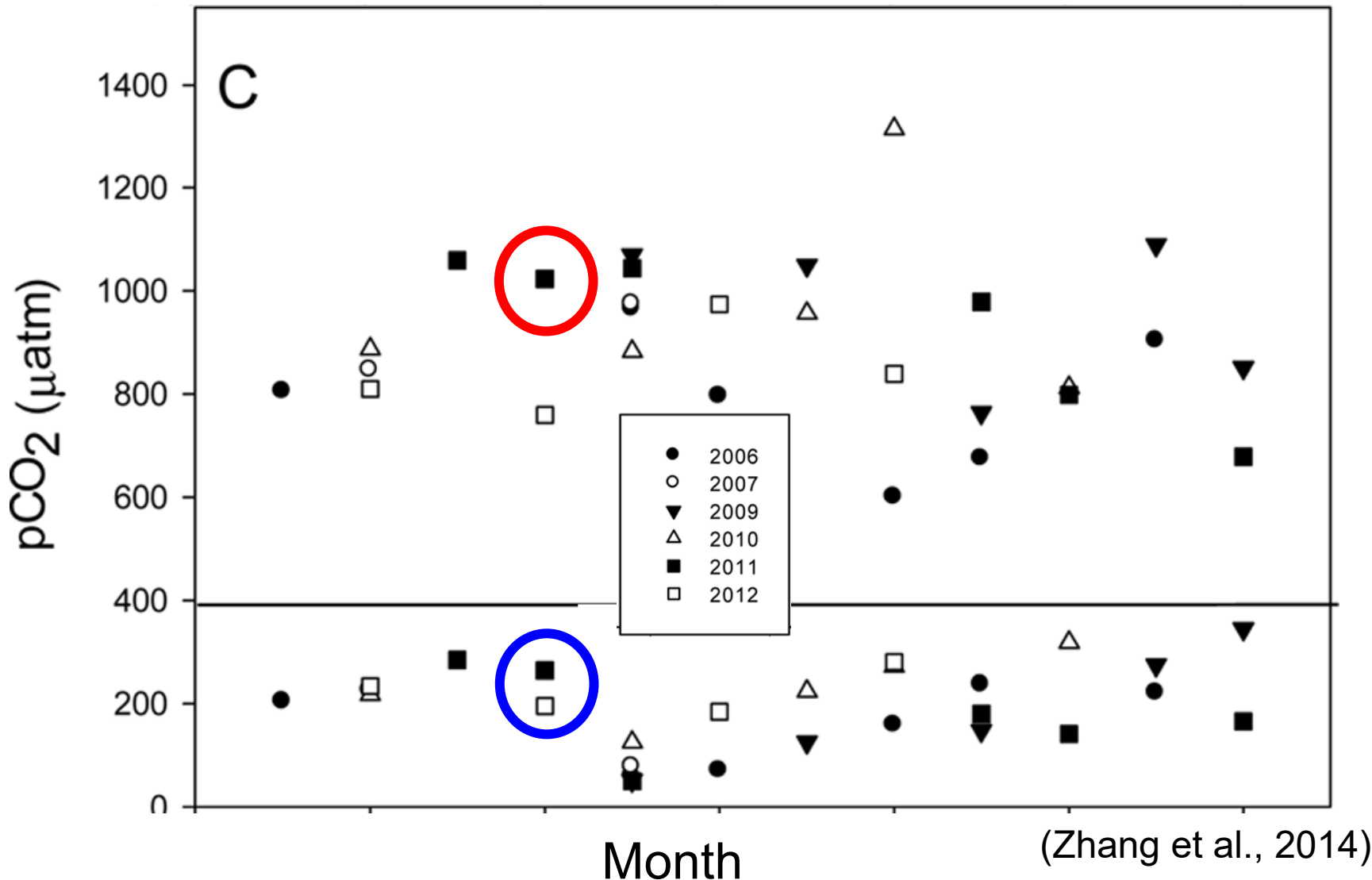
CO₂ flux near St. Bob Allen keys in 2019, from eddy covariance



Even in **Spring**, CO₂ from **sea** to **air** might be due to scale of bloom.

$6.1 \text{ mol m}^{-2} \text{ year}^{-1} > 1.50 \text{ mol m}^{-2} \text{ year}^{-1}$ in Zhang et al. 2014

Horizontally max and min pCO₂ in Florida bay

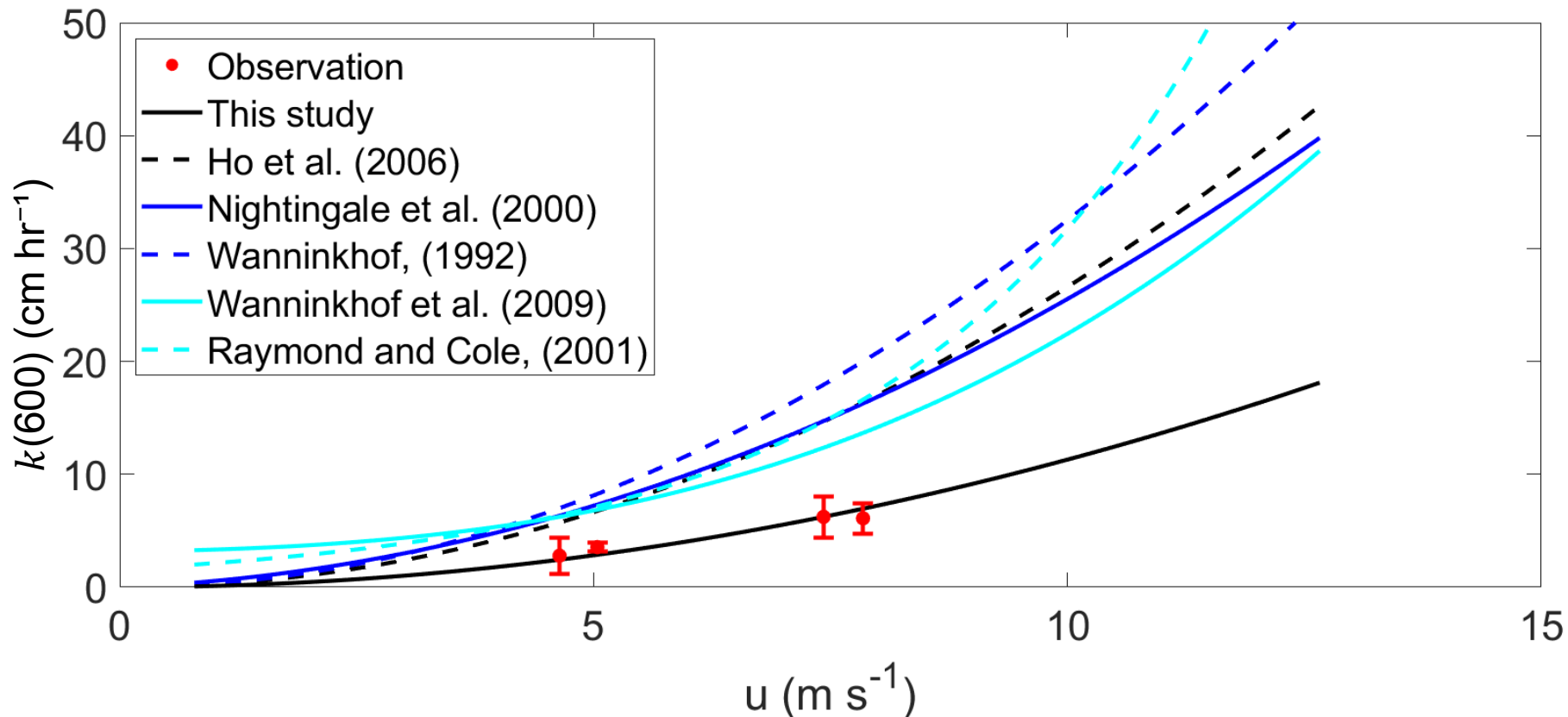


Min pCO_{2water} 200 ~ 1000 µatm

This study's result: pCO_{2water} = 228 µatm ← still in the range

Calculated $k(600)$

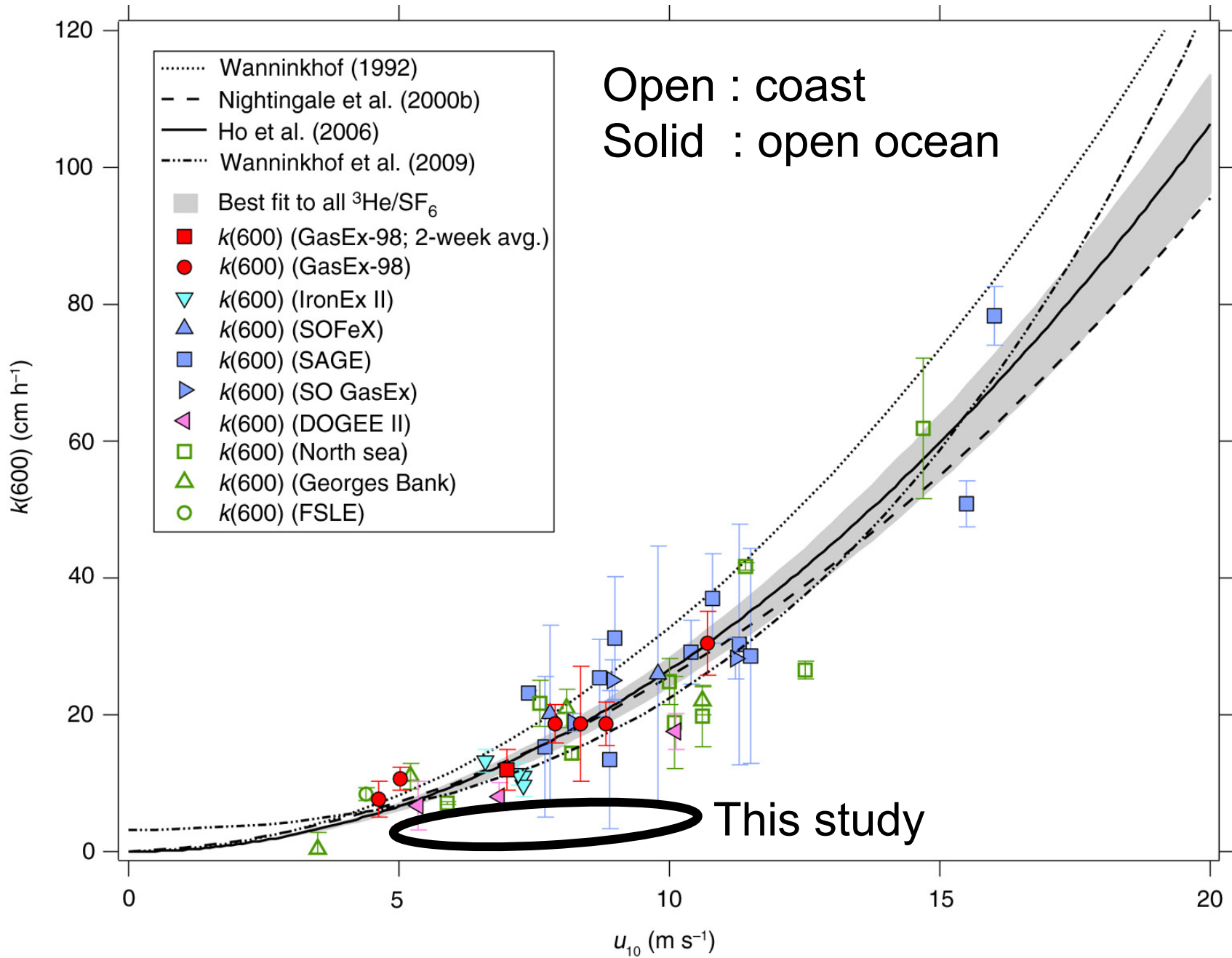
k for CO₂ at 20 °C



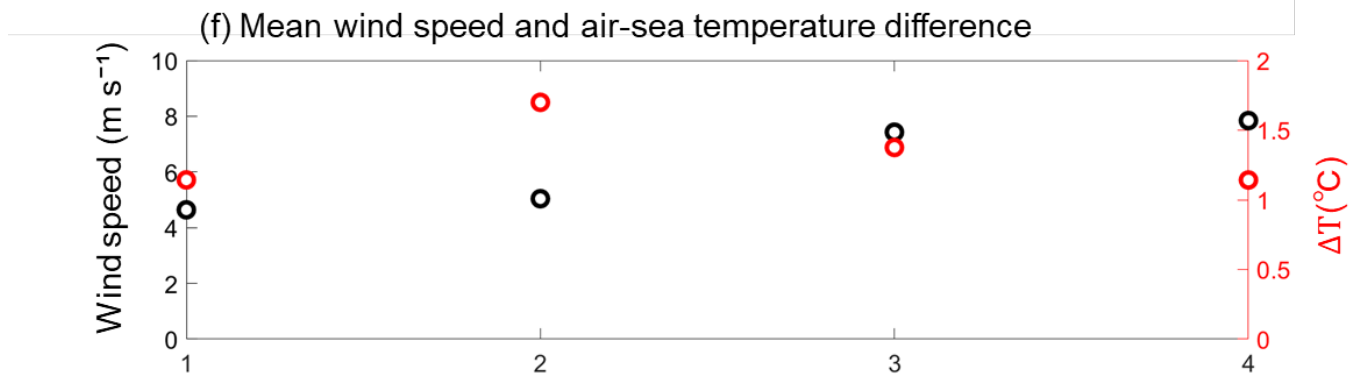
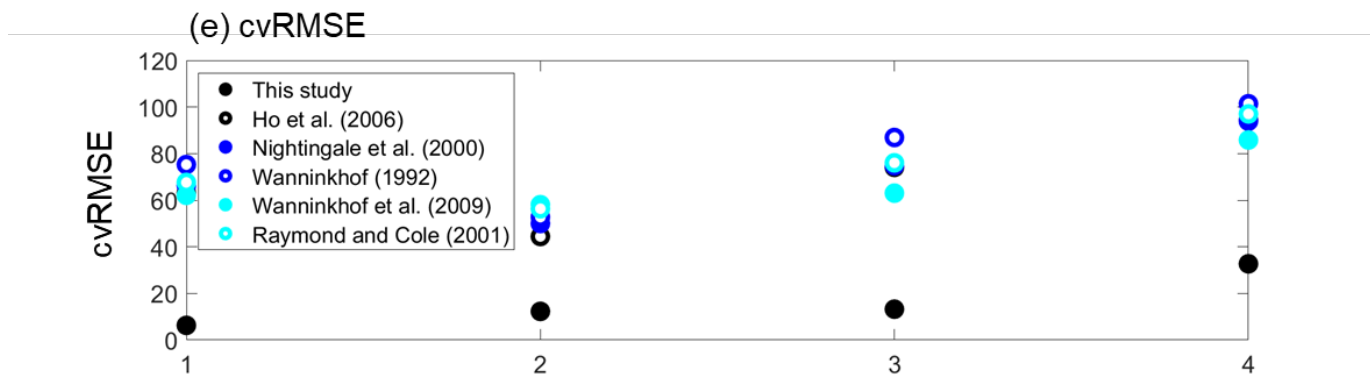
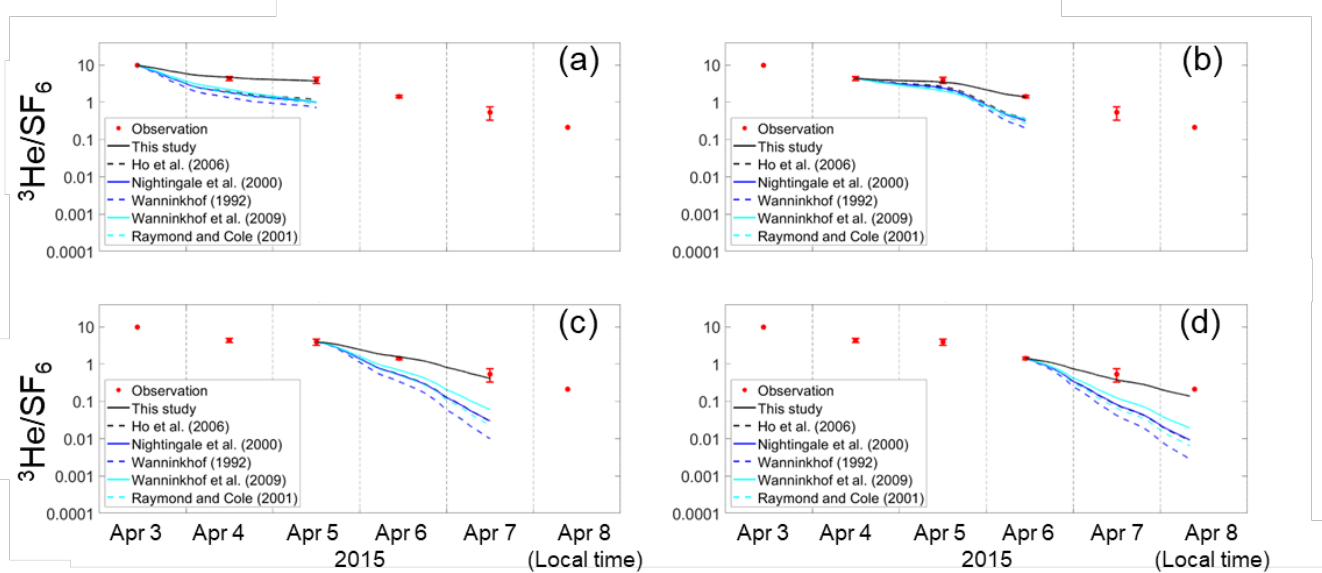
Average $k(600)$ is 4.67 ± 1.76 cm hr⁻¹

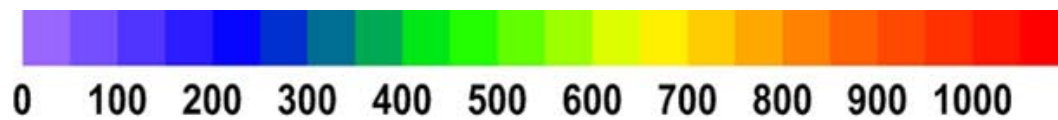
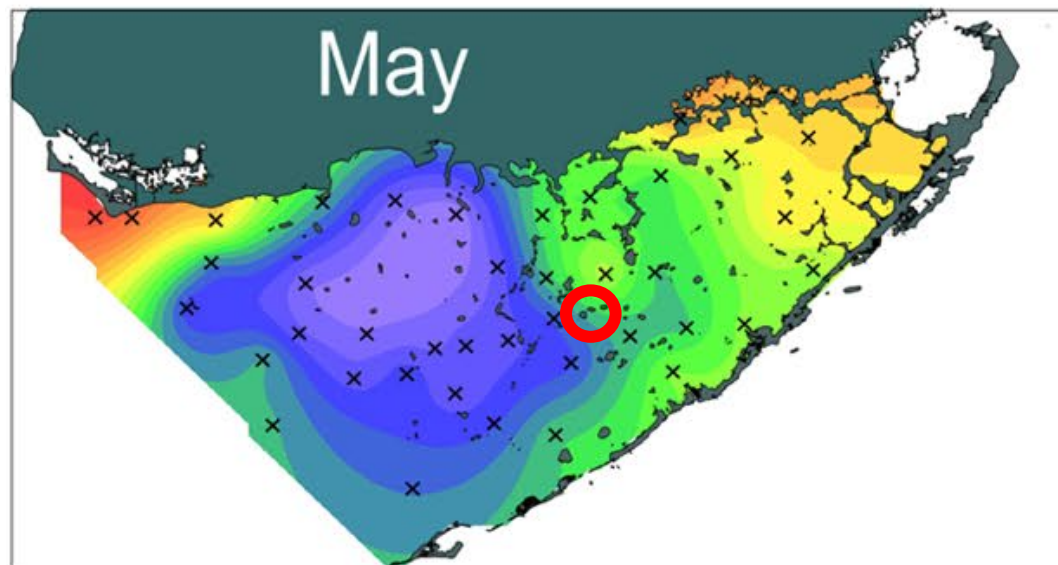
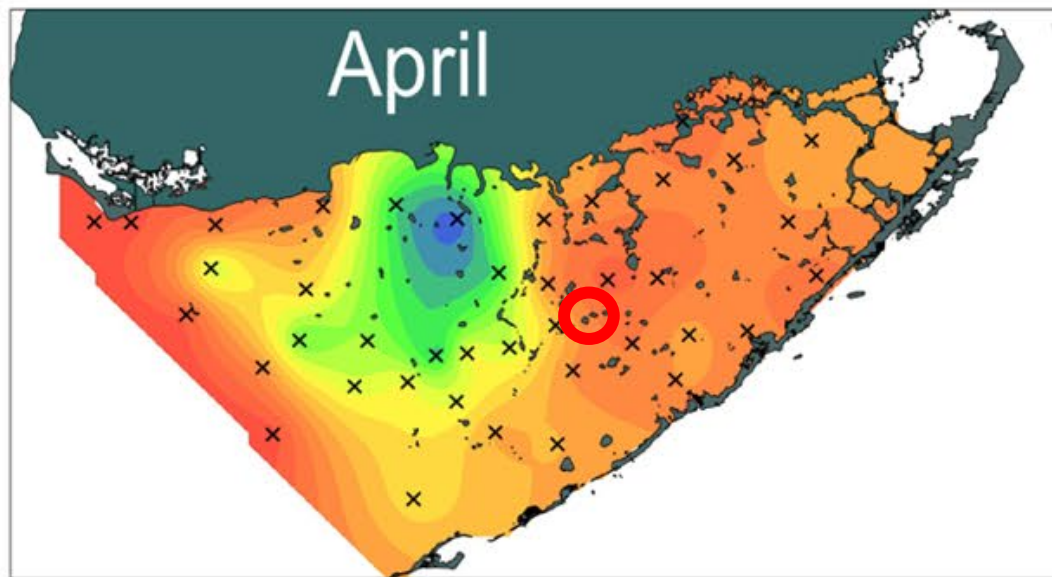
$k(600)$ is lower than previous experiments in coastal & open ocean

New parameterization, $k(600) = 0.122u_{10}^2$ fits well.



(Ho and Wanninkhof, 2016)

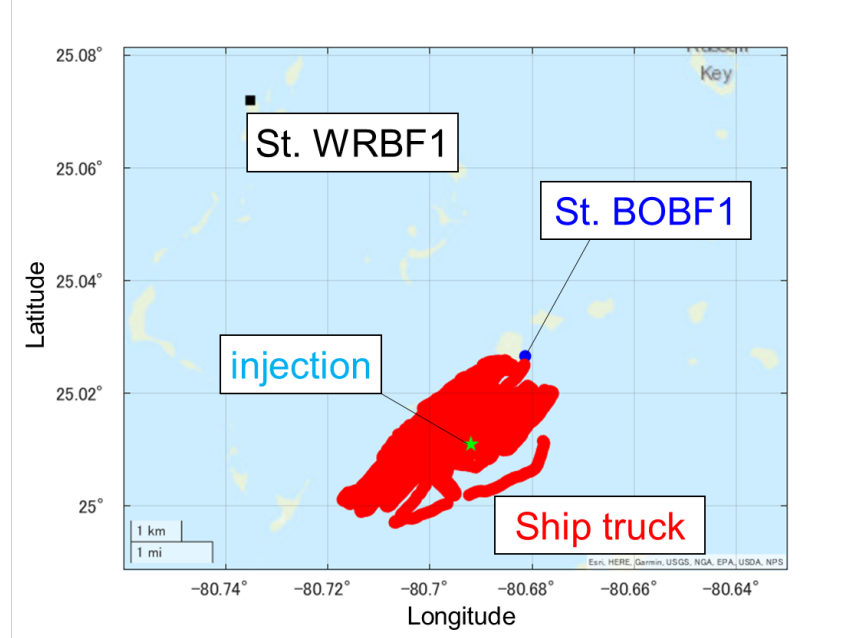




pCO₂(μatm)

(Zhang et al., 2014)

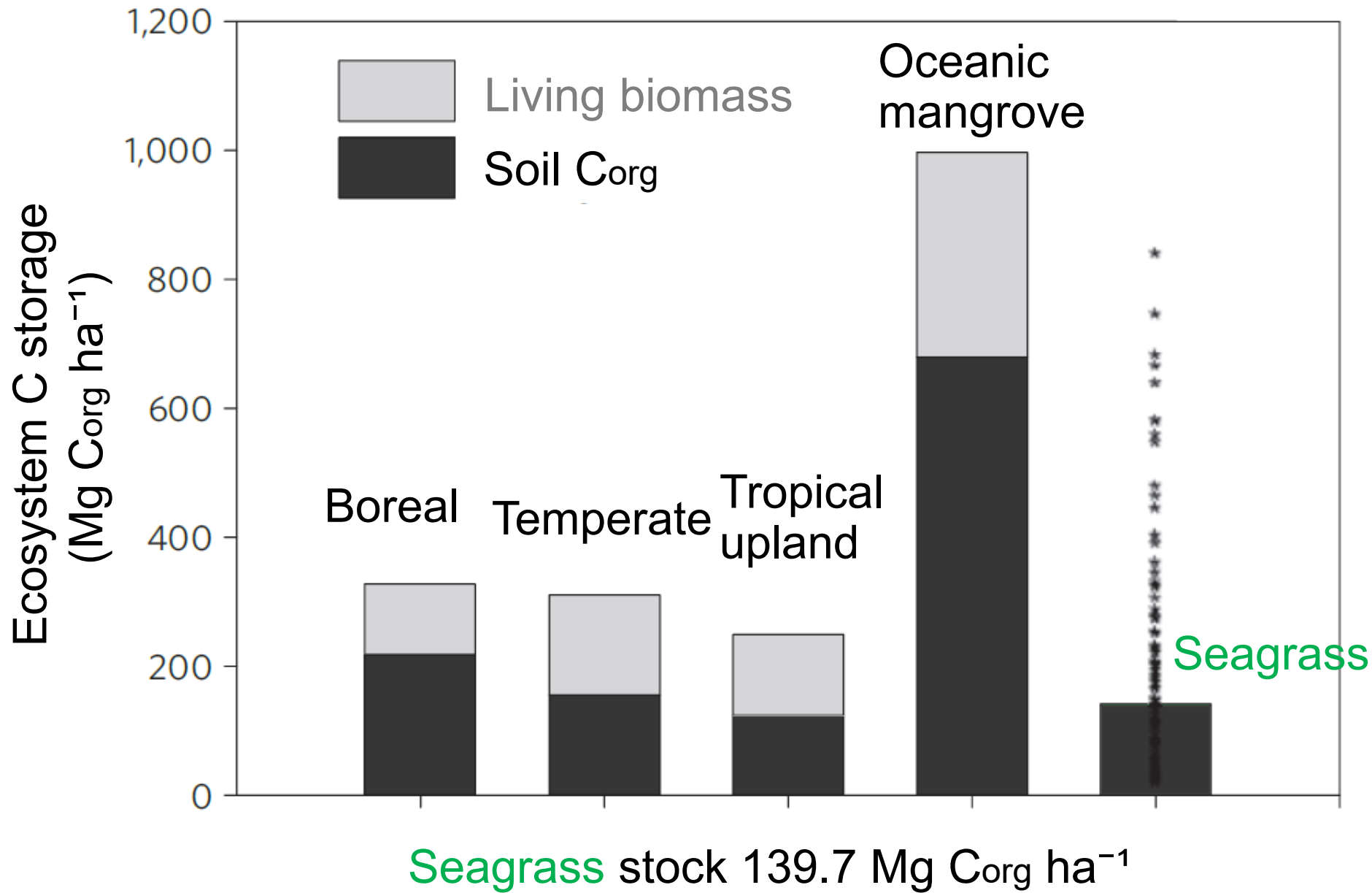
Methods



St. BOBF1 & WRBF 1

parameters	Method
Wind	Sonic anemometer. Converted to 10 m height
Air T	Sonic anemometer
Water T	Taken by Everglades National Park

Seagrass's role in Global carbon cycle



(Fourqurean et al., 2012)